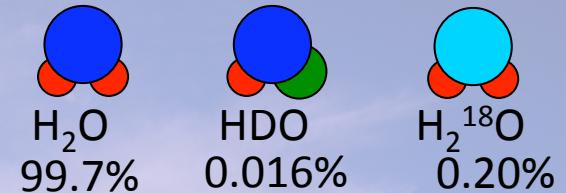


Diagnosis of the present hydrological cycle using isotope GCM

Naoyuki Kurita
(Visiting Scientist from JAMSTEC)

Outline

1. New water isotopes dataset in Asia
2. Findings of new dataset
3. Factors controlling the isotopic variation using isotope GCM
4. How to use the isotope tracer to diagnosis of the hydrological cycle



Borneo island

Indian Ocean

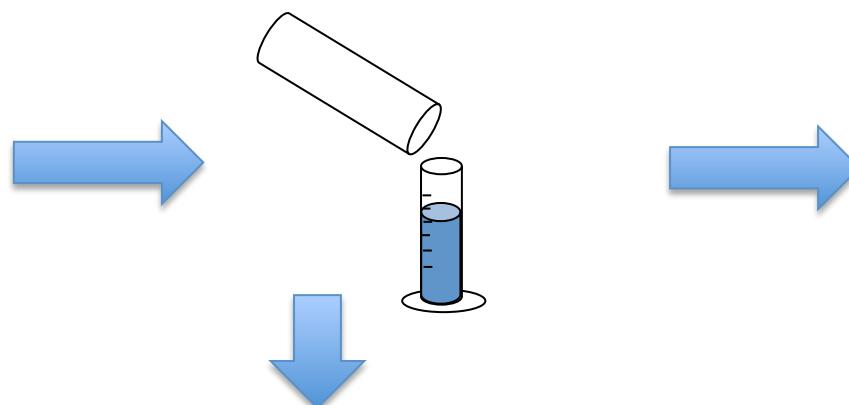
Water Isotope Sampling Procedure

1. Collect rain samples

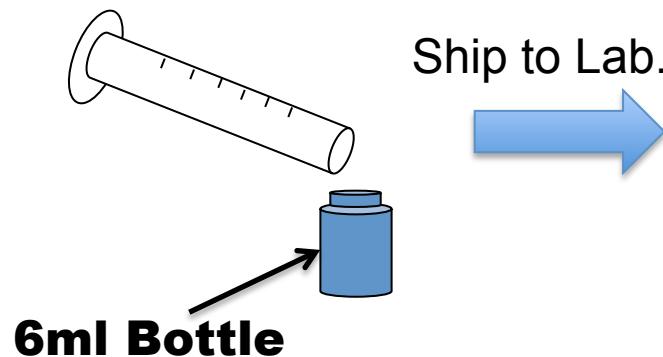


Russian station

2. Measure the amount of precipitation



3. Replace rain water to glass bottle (6ml)



3. Send to WMO through the telecommunication network system

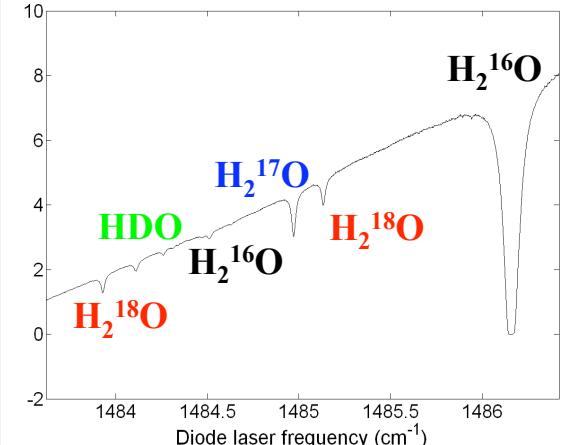
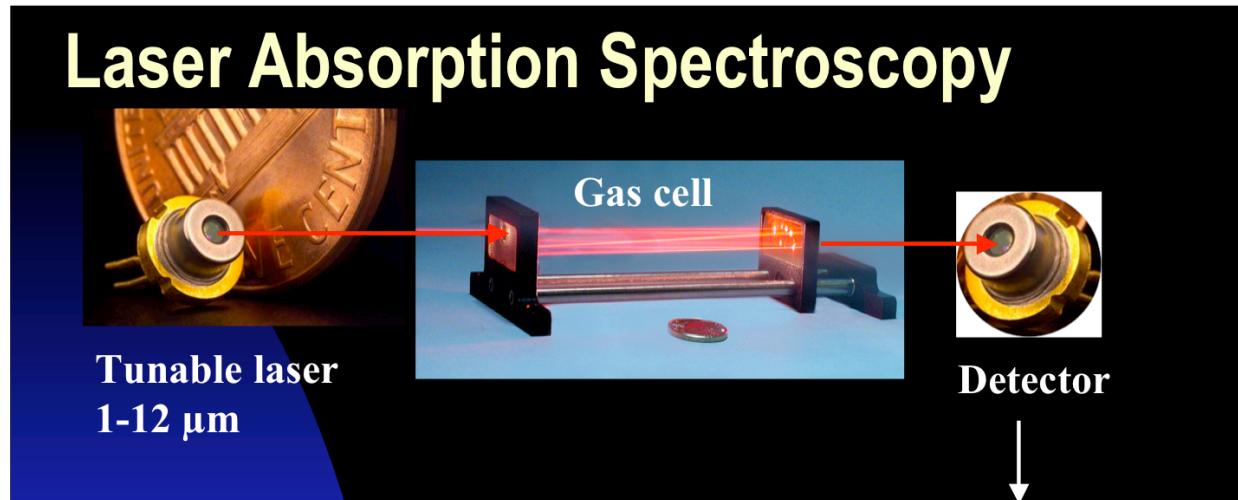


World Meteorological Organization

4. Measure the isotopic content of prec.



New Measuring Technique



In-situ observation

Airplane



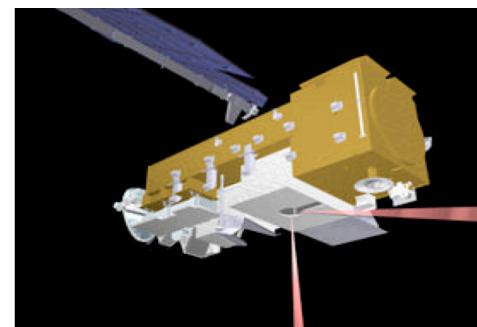
NASA WB-57

Balloon



TETHEO campaign

Satellite



AURA/TES

New Measuring Technique

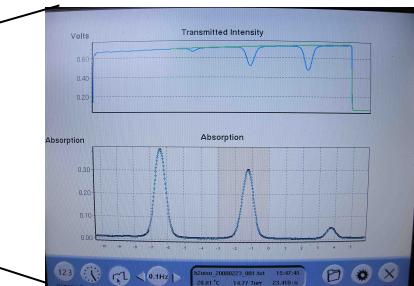
Laboratory → Reduce the measuring time



CF-IRMS

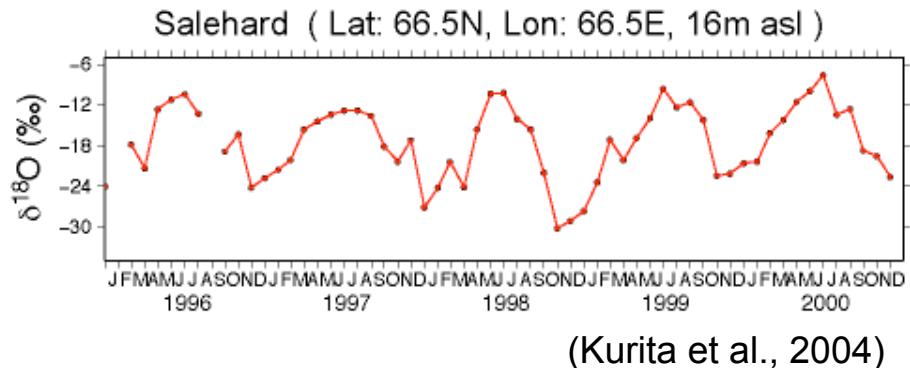


Laser mass spec.

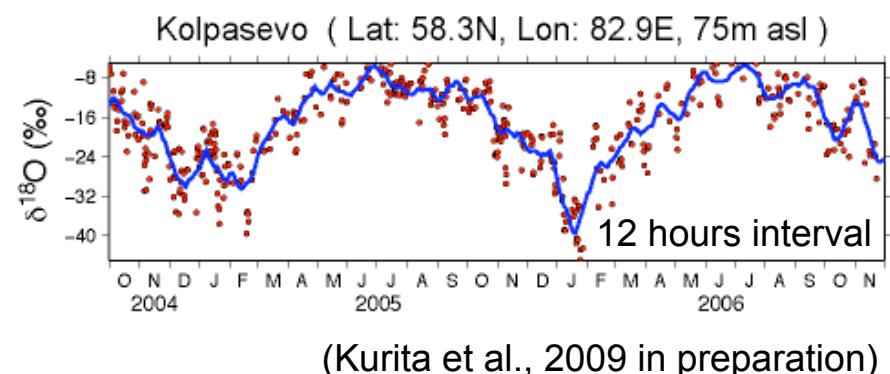


Monitoring datasets

Monthly archive → Event sampling

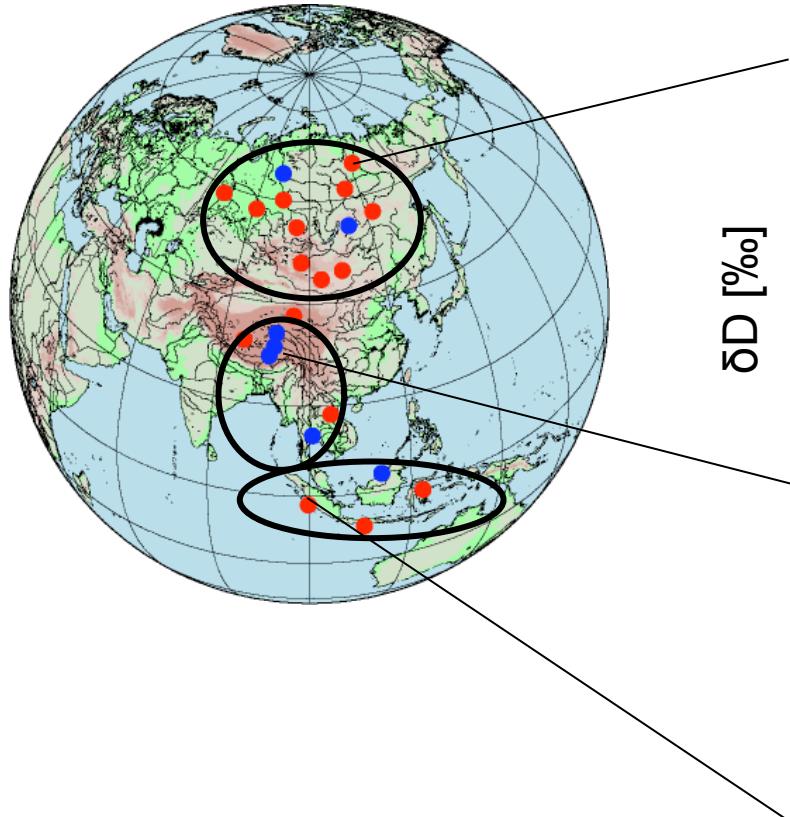


(Kurita et al., 2004)

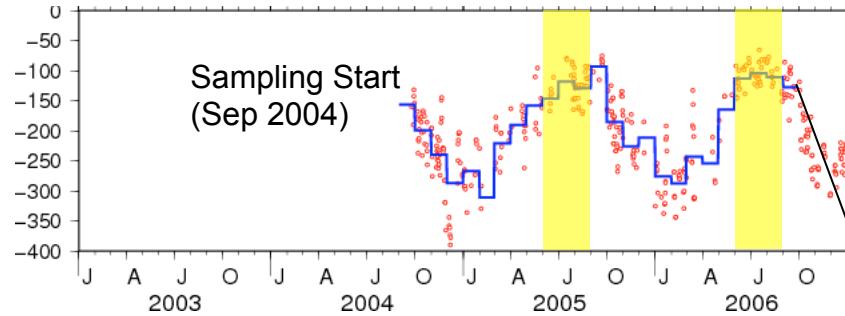


(Kurita et al., 2009 in preparation)

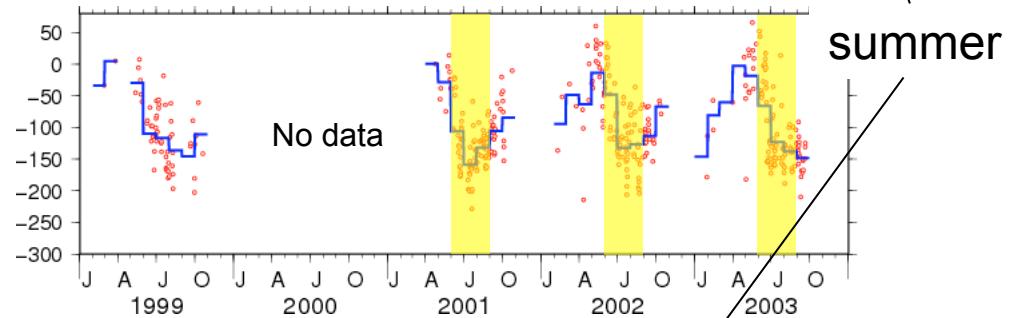
Water Isotope Monitoring in Asia



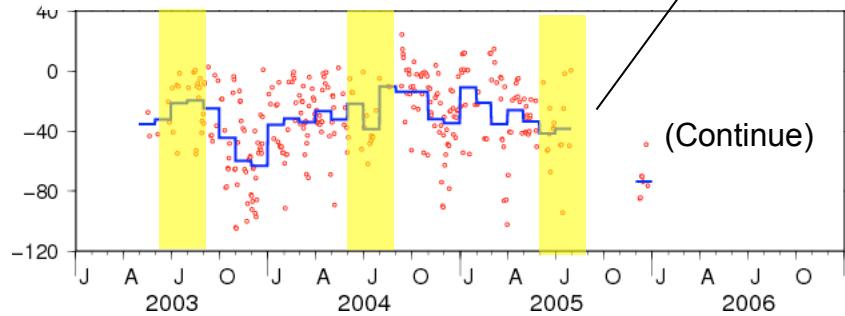
Siberia (Zhigansk, 66.77N, 123.42E)



Tibet (Lhasa, 29.67N, 91.13E)

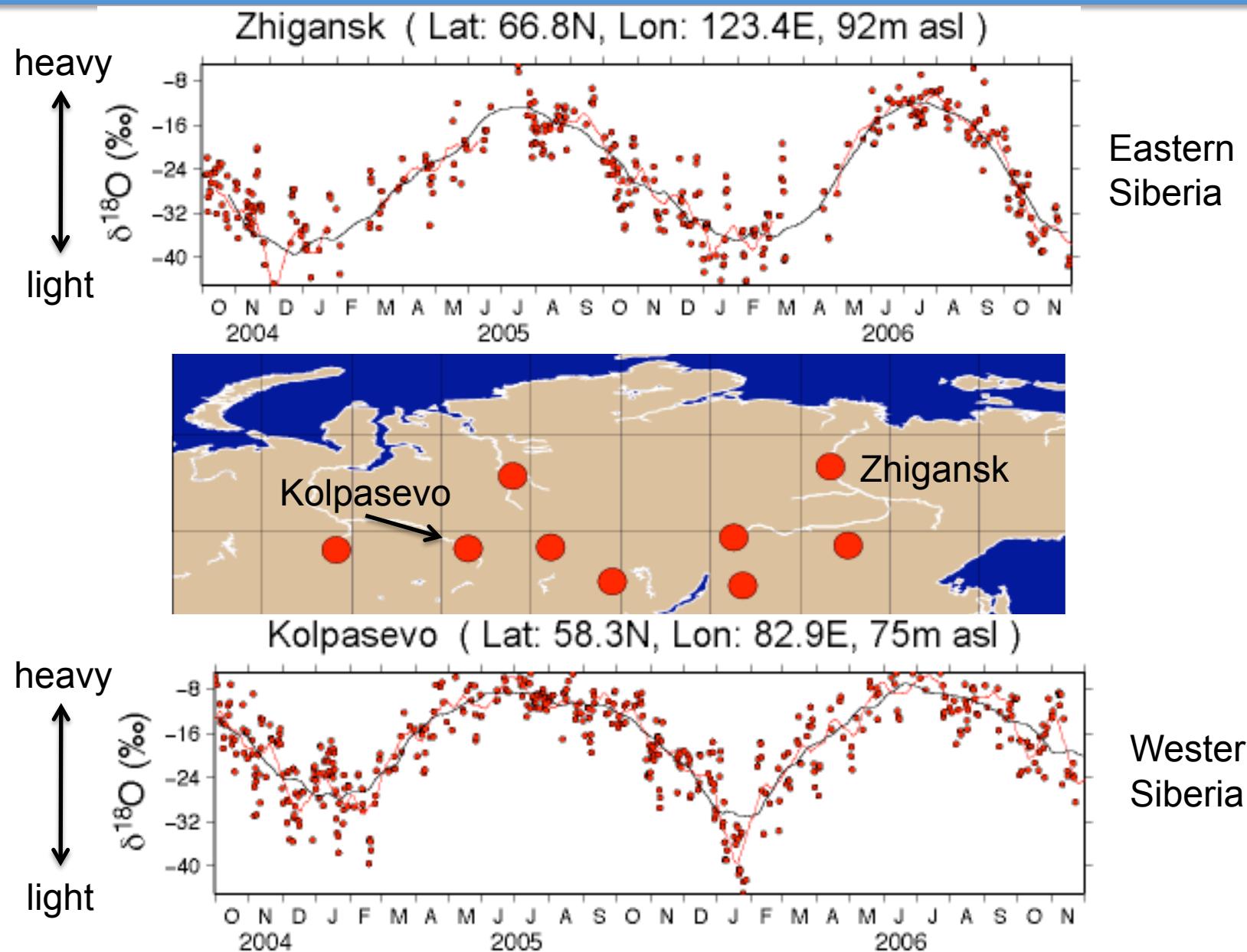


Maritime Cont. (Jambi, 1.63S, 103.65E)



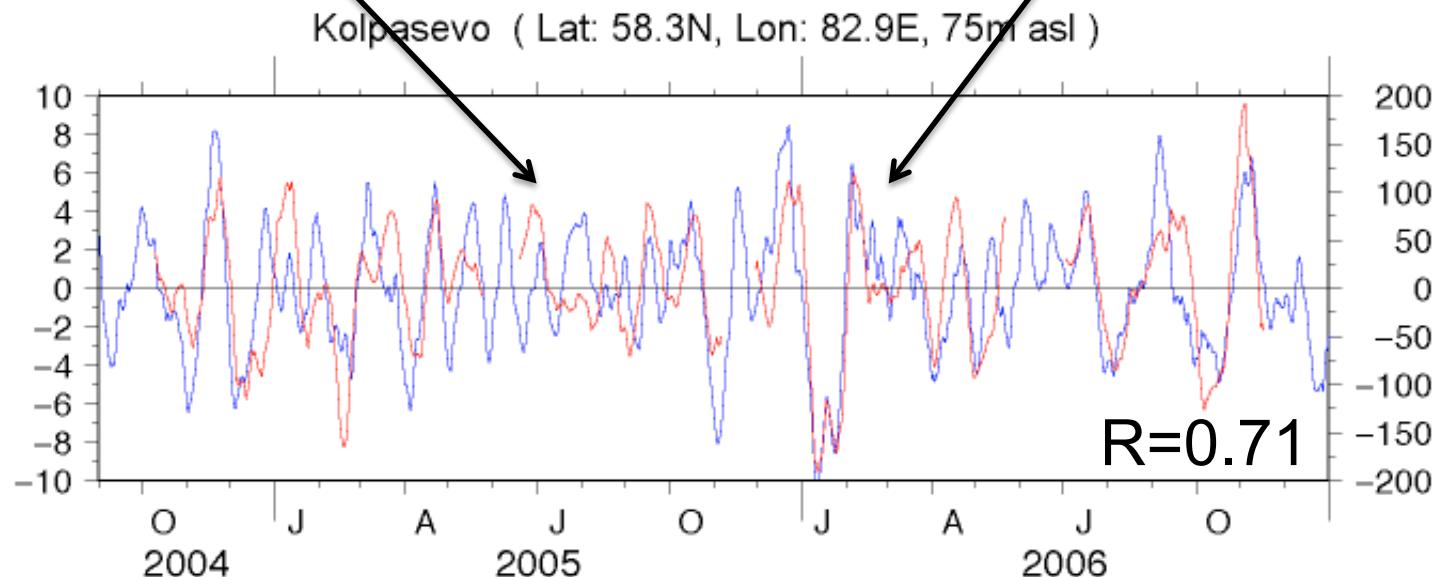
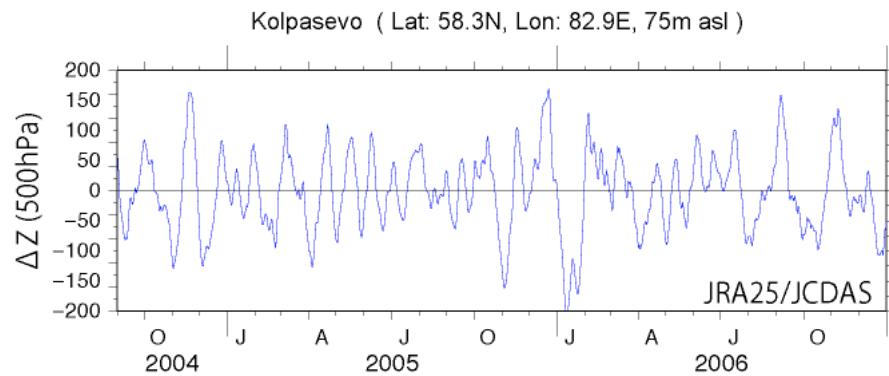
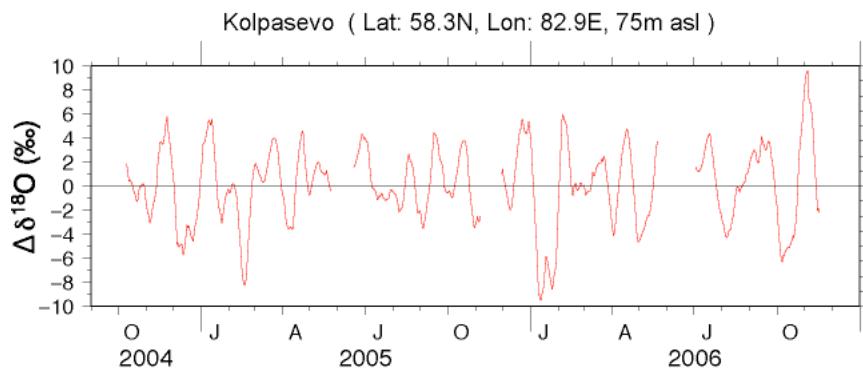
What can we get from these high frequency monitoring datasets?

Time series of $\delta^{18}\text{O}$ in precipitation in Siberia

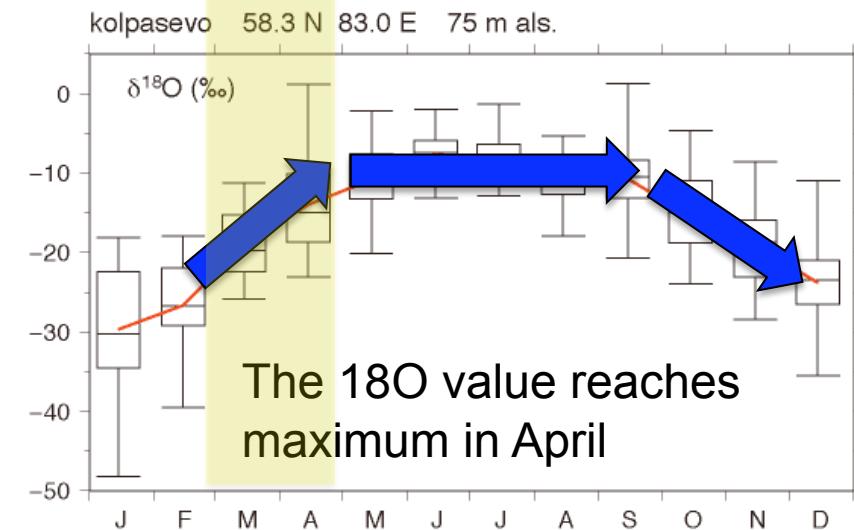
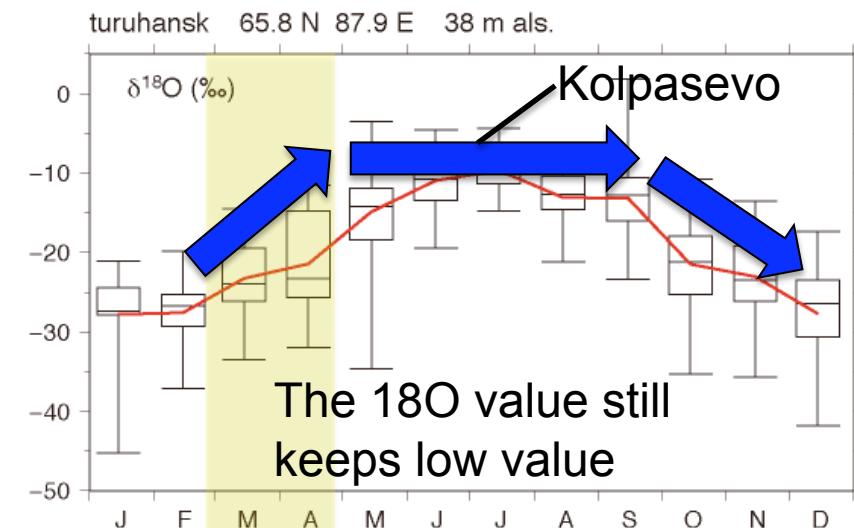
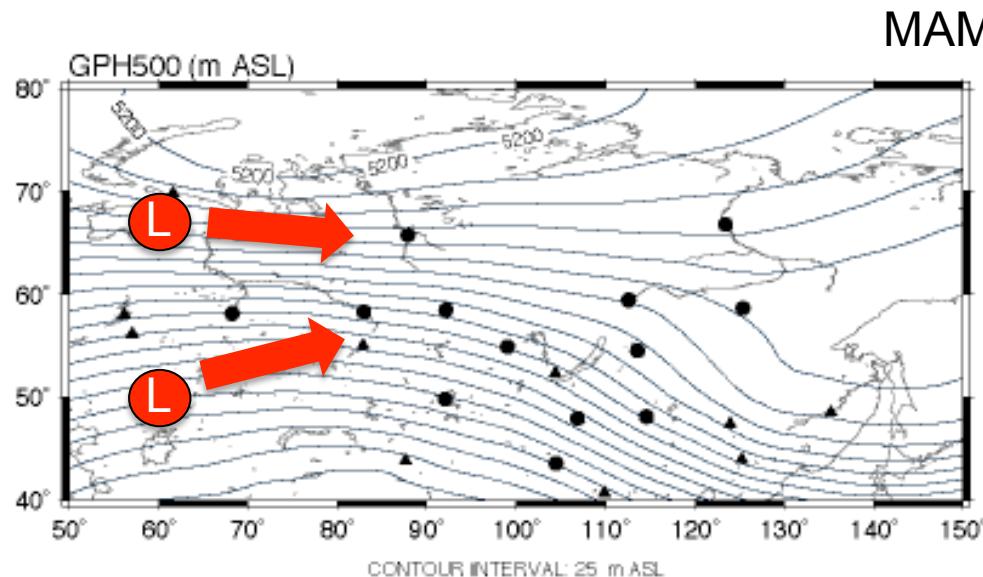
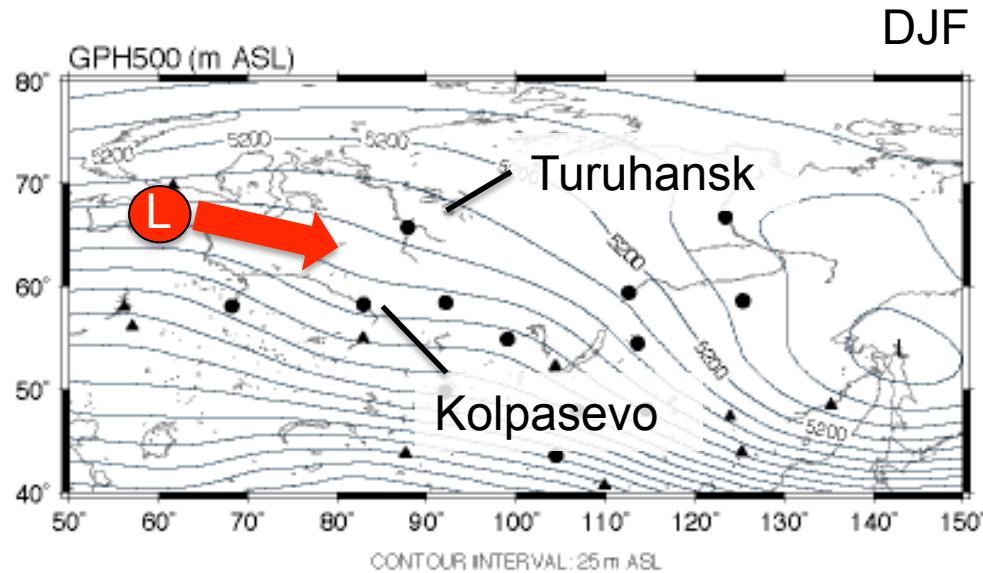


Intra-seasonal variability of $\delta^{18}\text{O}_{\text{pre}}$ in Siberia

10-30 day band passed filtered anomaly

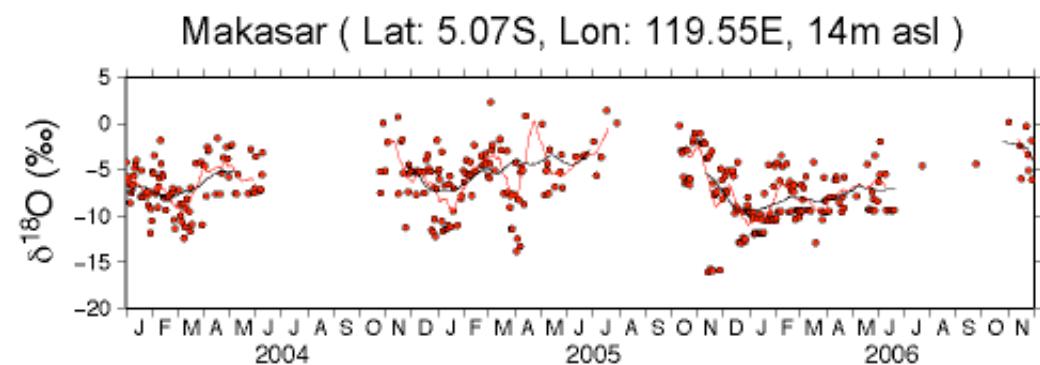
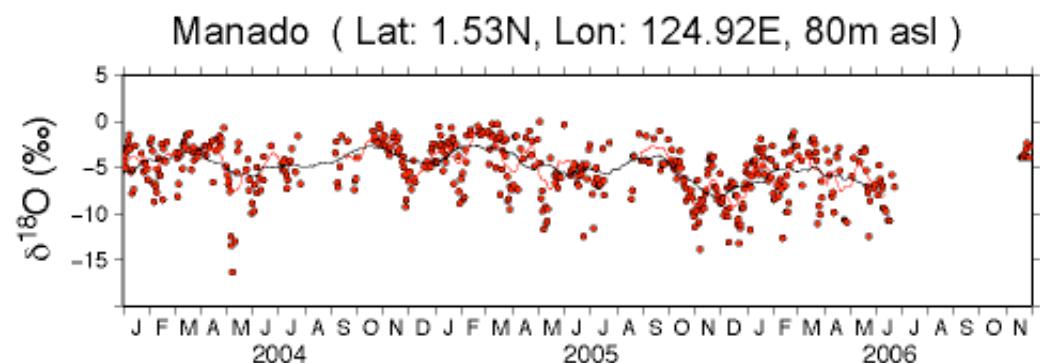
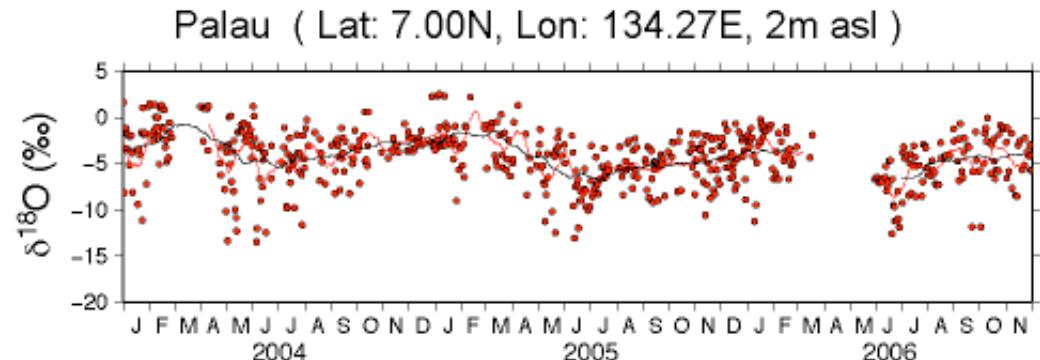
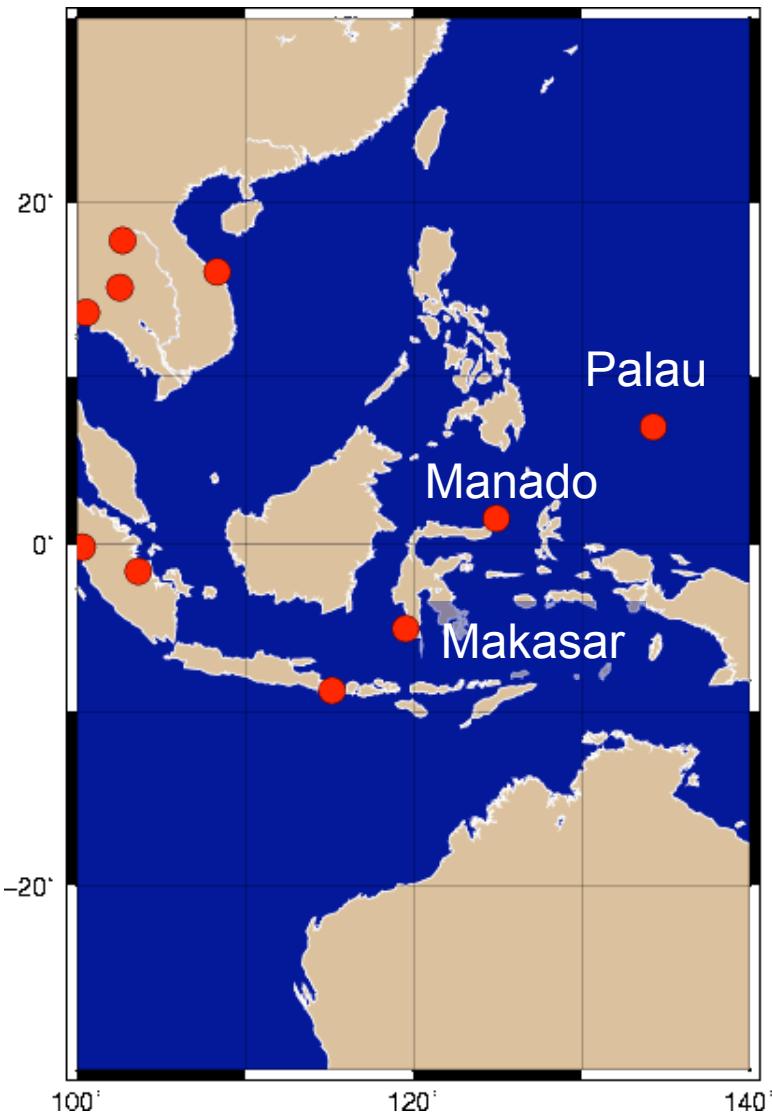


Seasonal cycle of $\delta^{18}\text{O}_{\text{pre}}$ in Siberia



Time series of $\delta^{18}\text{O}$ in precipitation in Tropics

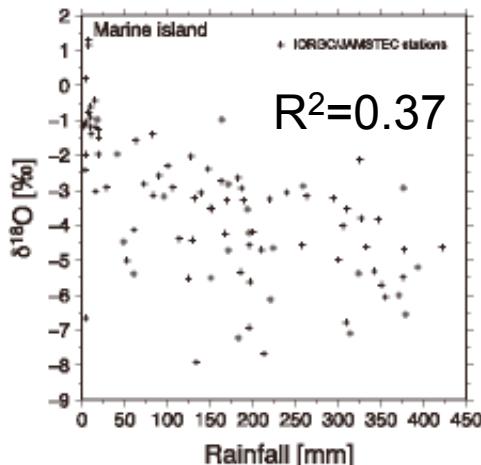
Intra-seasonal $\delta^{18}\text{O}$ variation >> Seasonal $\delta^{18}\text{O}$ variation



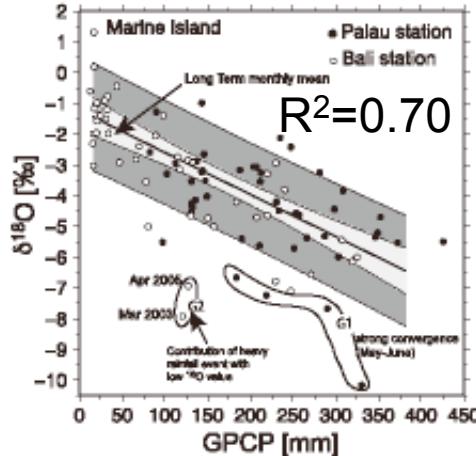
Isotope/Rainfall amount relationship (monthly)

Sub-tropical region

Local Rainfall



Regional Rainfall

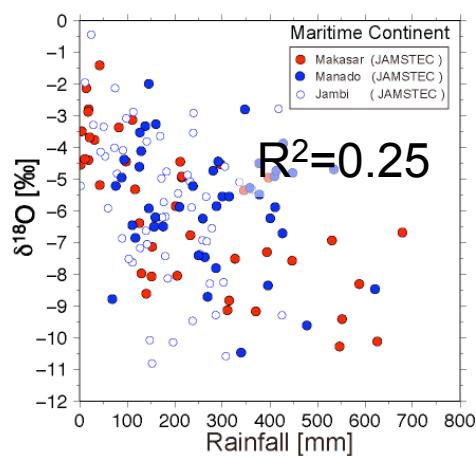


On a monthly timescale, the correlation of $\delta^{18}\text{O}$ value with regionally averaged precipitation amount was much higher than that with station based precipitation.

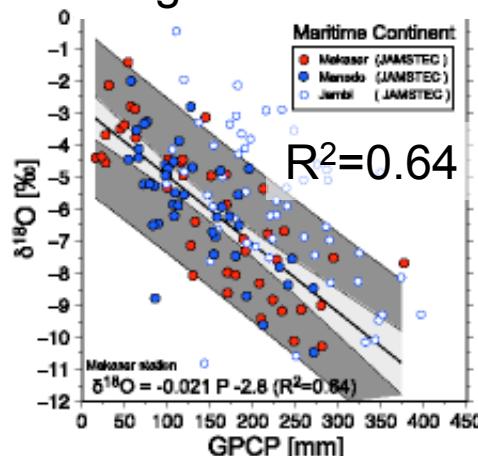


Maritime Continent region

Local Rainfall



Regional Rainfall



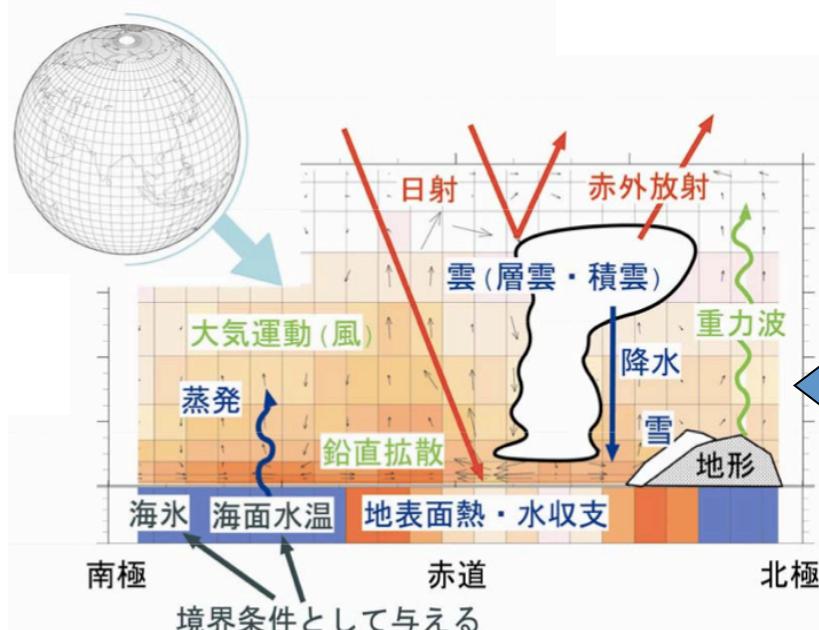
The variability of $\delta^{18}\text{O}$ values observed at a station is related with the regional scale hydrological circulation.

(Kurita et al., 2009)

Water Isotope GCM

Isotope scheme is incorporated into the physics scheme of GCM

MIROC ver3.2(IPCC AR4)



Isotope physics

Isotope changes only occurs at condensation and surface flux scheme

Convective rain



Evaporation



Stratiformed rain



snow



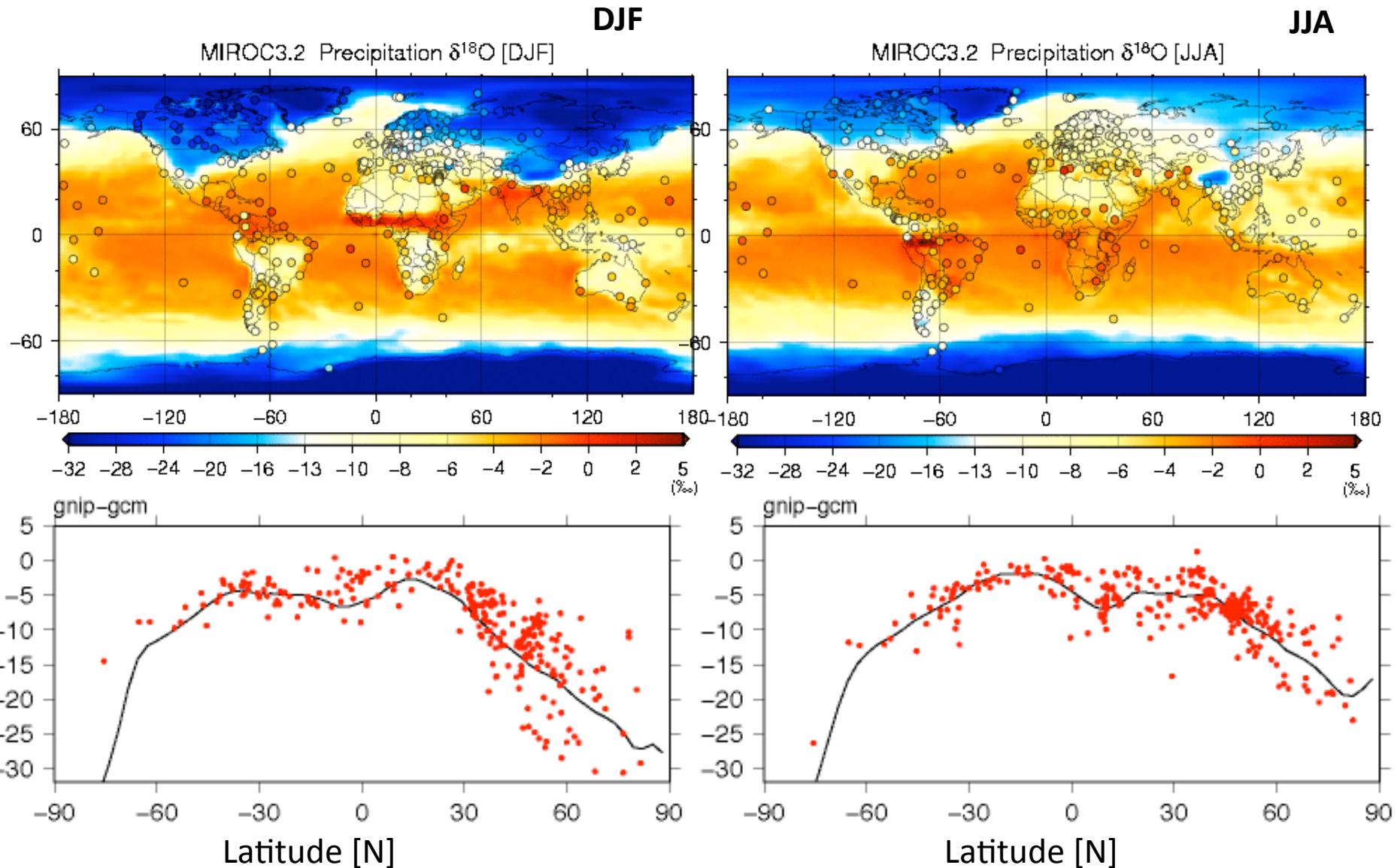
陸上

Definition of Isotope value

$$\delta^{18}\text{O} = \left(\frac{Q(H_2^{18}\text{O})}{Q(H_2\text{O})} - 1 \right) \times 1000$$

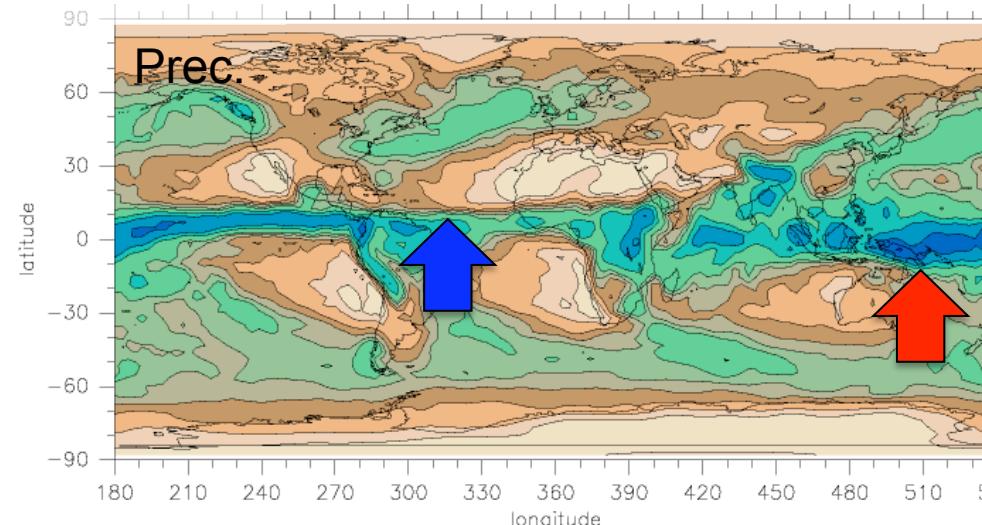
Q(TH₂O):Total moisture
(vapor+cloud liquid)

Model performance (H_2^{18}O)

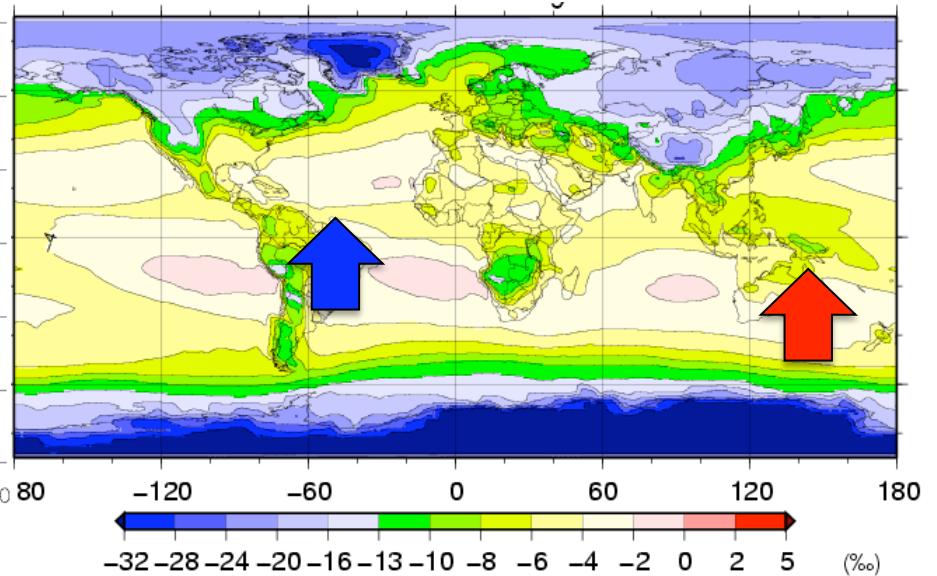
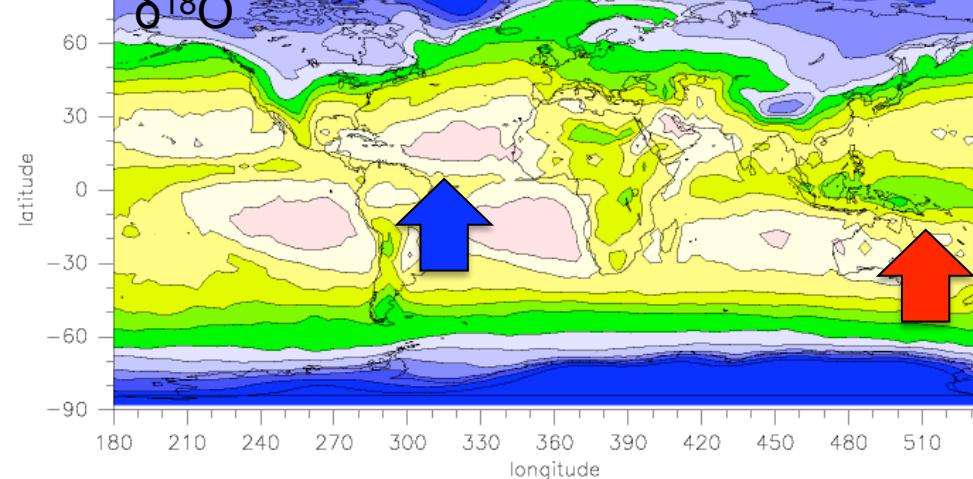
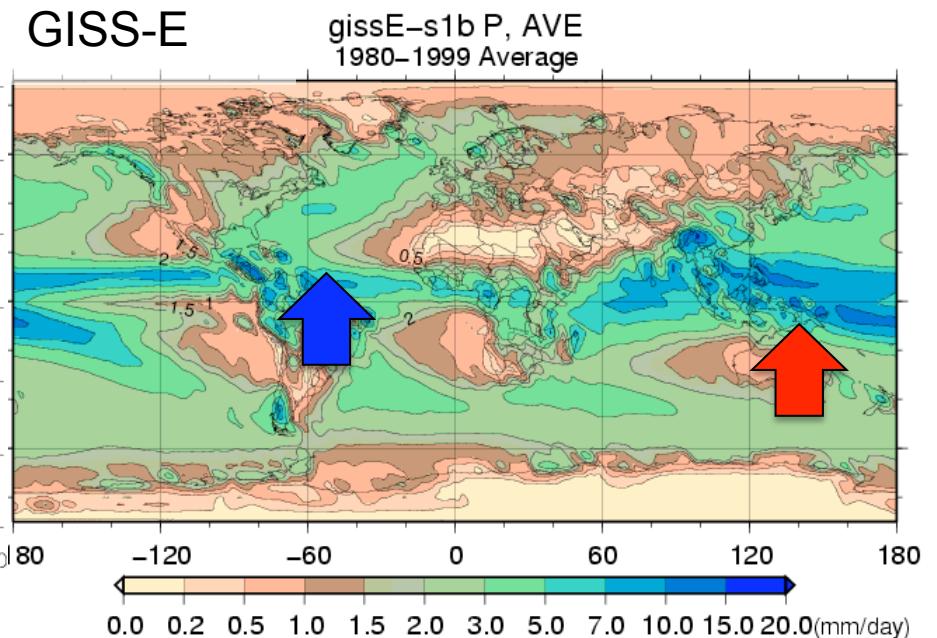


Comparison between GISS-E and MIROC

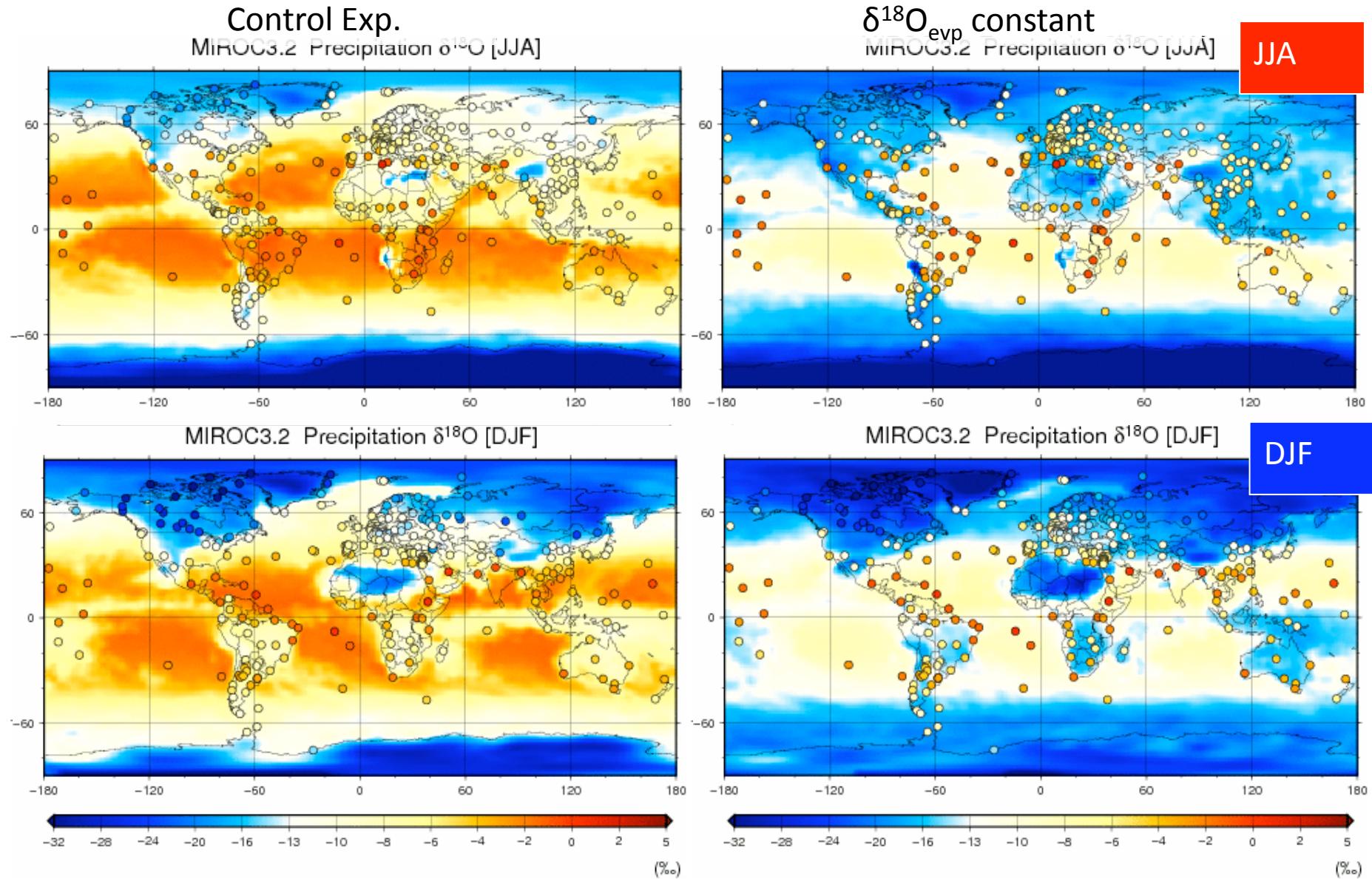
MIROC



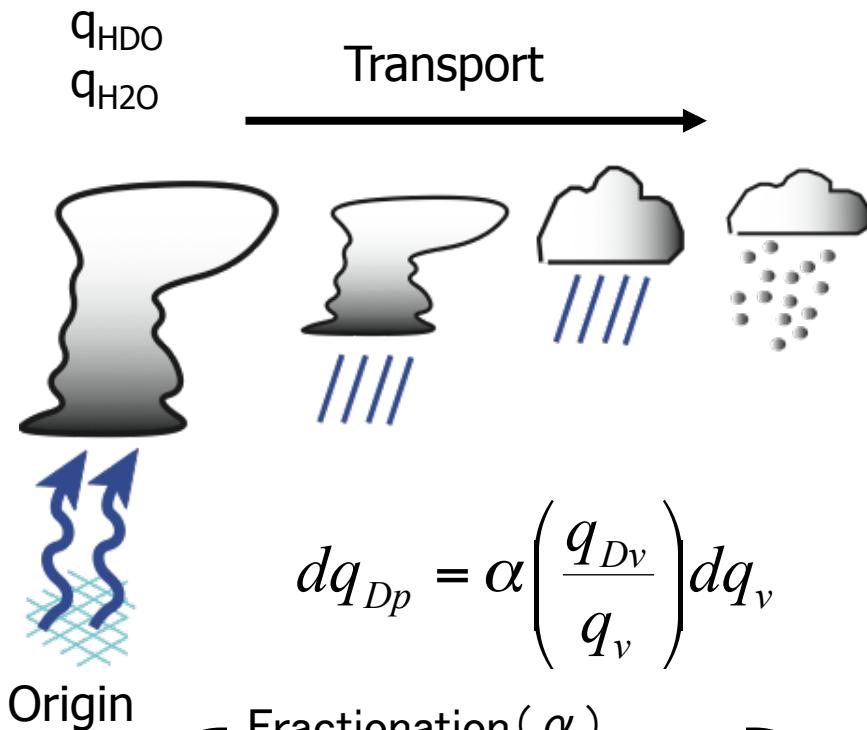
GISS-E



Sensitivity Test (Source effect)



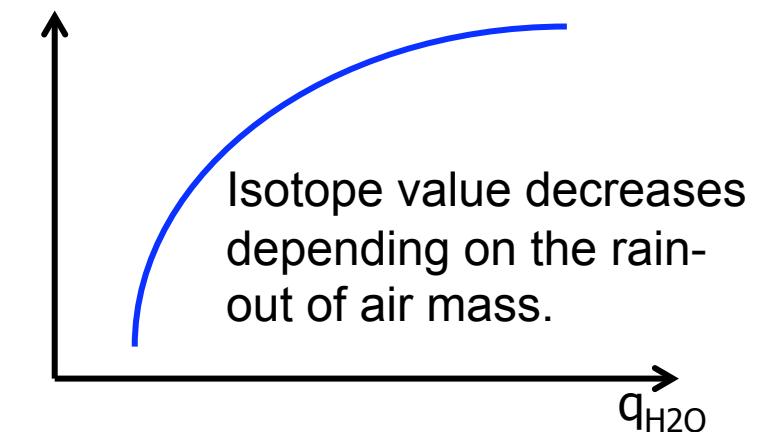
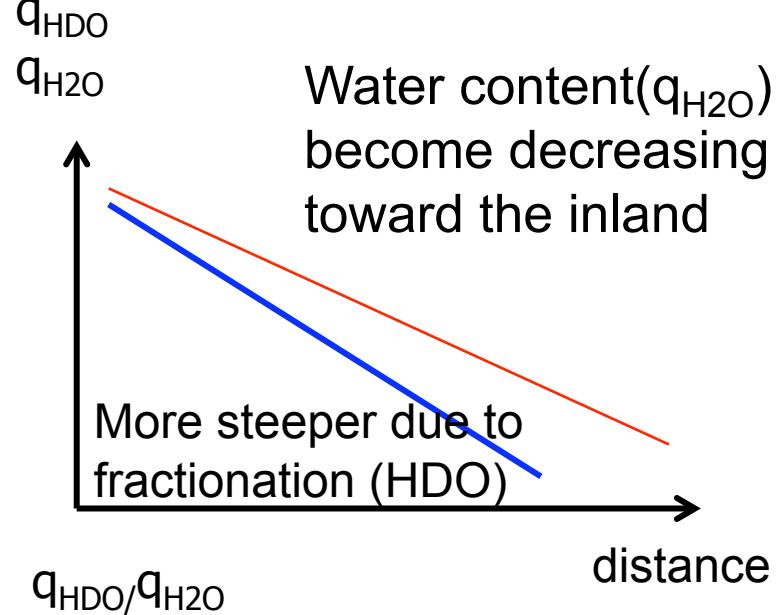
Rain-out effect



Origin

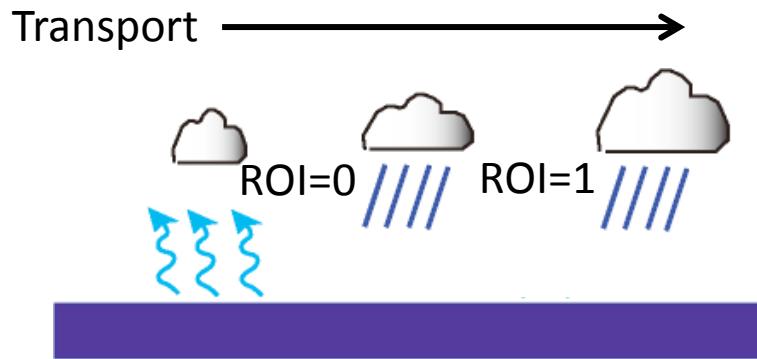
Fractionation (α)
heavy isotopes tend to
enriched in condensation
phase

$$\delta D \text{ (liquid)} > \alpha \delta D \text{ (vapor)} \quad \alpha > 1$$



Rain-Out Index (ROI)

What is the ROI?



The concentration is set to zero at the time of evaporation from the ocean, and then is forced to increase the amount of precipitation at each time step.

Basic Equation for water vapor

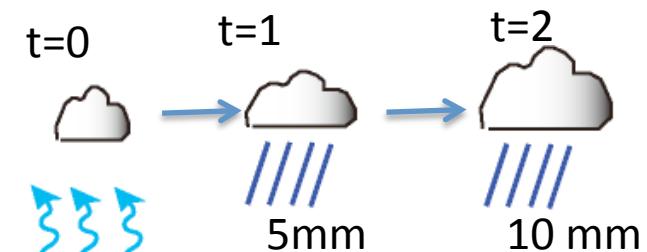
$$\frac{\partial q^t}{\partial t} = -v \nabla q^t - \frac{1}{\rho} \frac{\partial}{\partial z} (F^D + F^C - F^P)$$

The Equation for tracing water vapor

$$\frac{\partial}{\partial t} (xq^t) = -v \nabla (xq^t) - \frac{1}{\rho} \frac{\partial}{\partial z} (F_x^D + F_x^C - F_x^P) + q_P^t$$

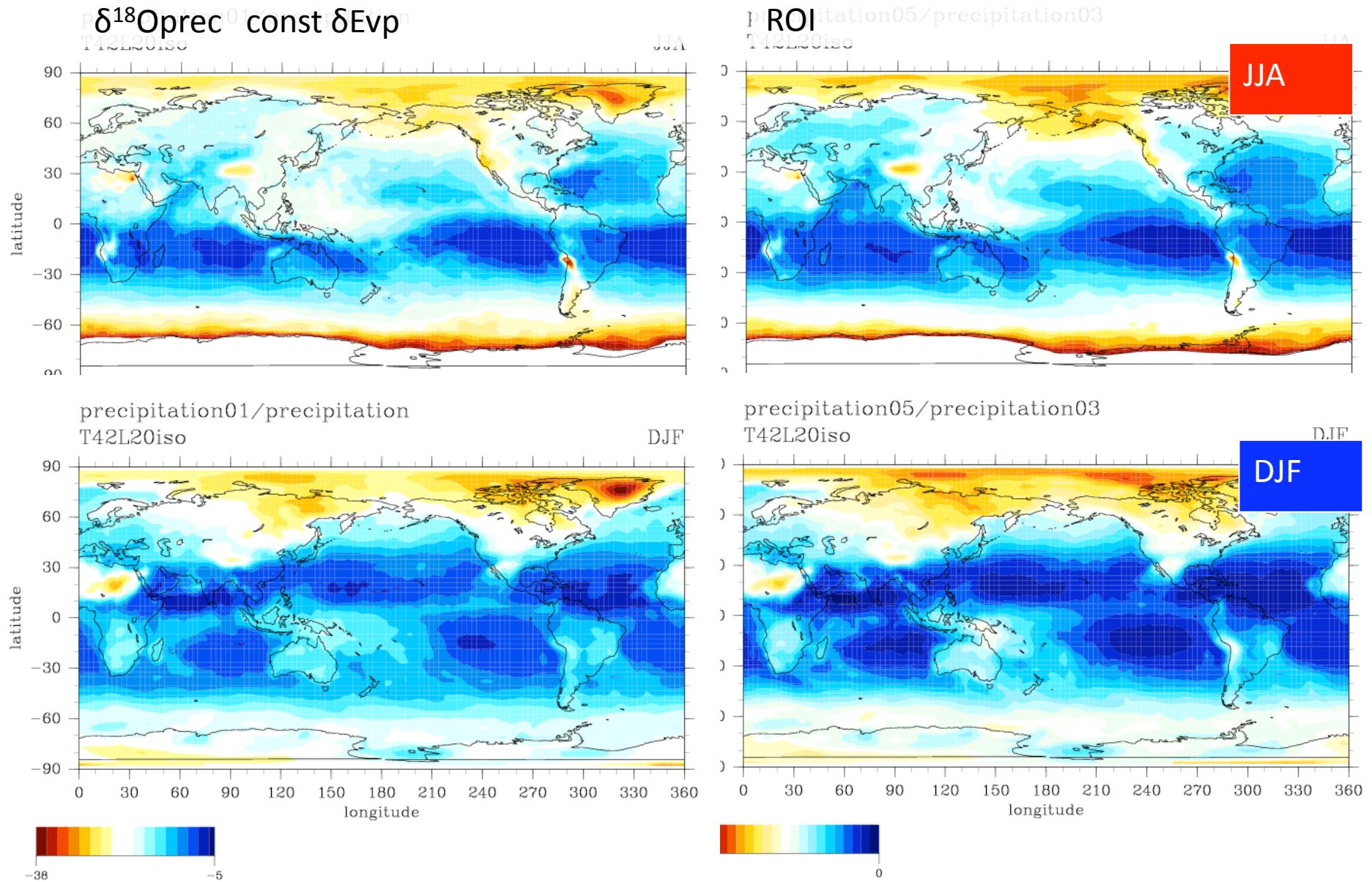
$$ROI = P_{xq}^t / P_q^t$$

For Example

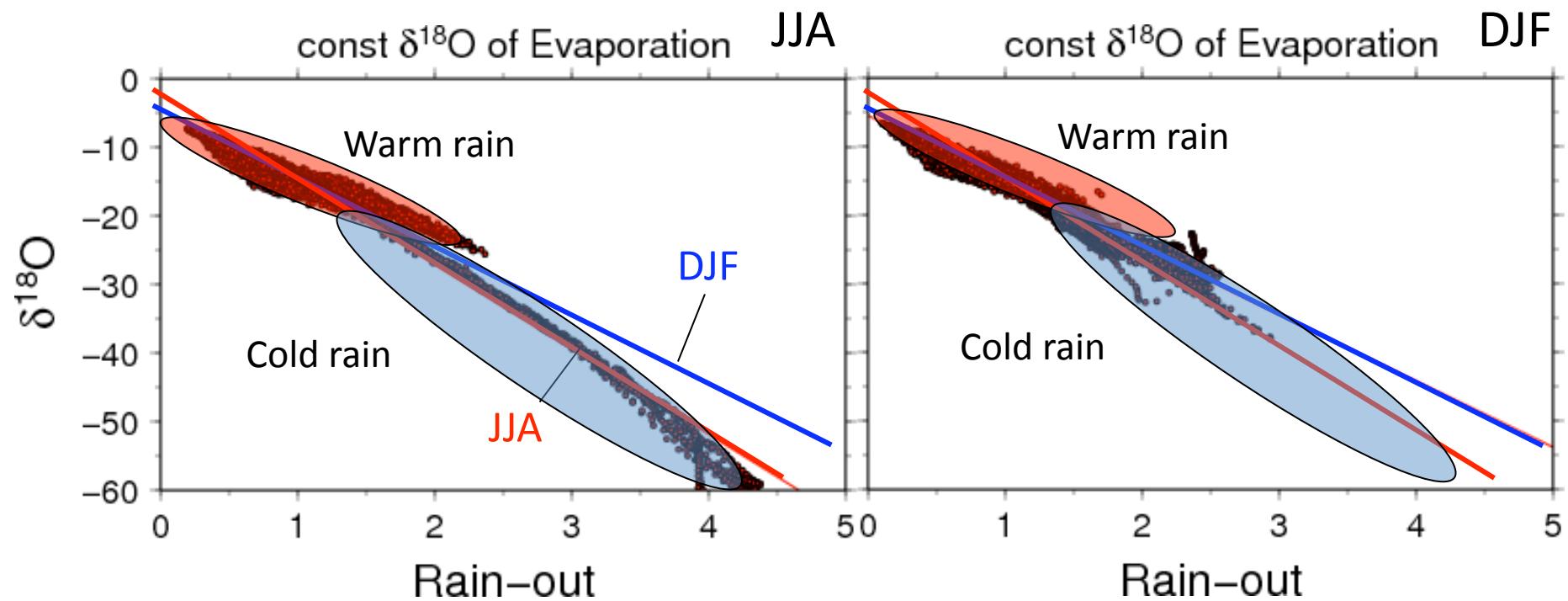


$$P_1 / P_2 = 5/10 \rightarrow ROI=0.5$$

Rain-out effect



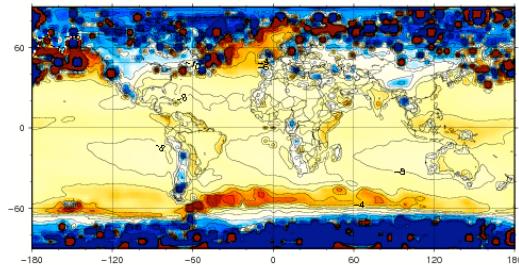
ROI vs $\delta^{18}\text{O}$ (const. δE)



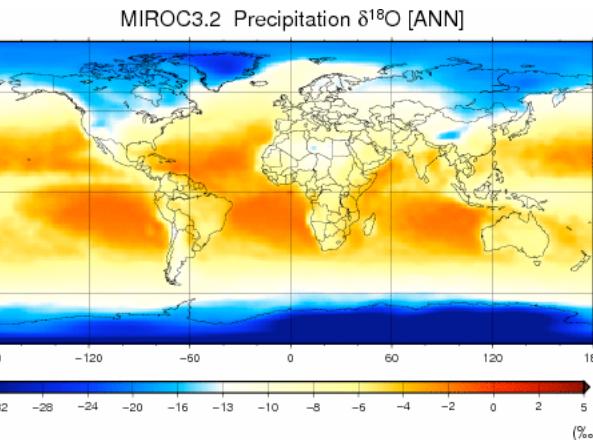
Global isotopic pattern seems to be controlled by the rain-out history from the source region.

Factors controlling isotopic value

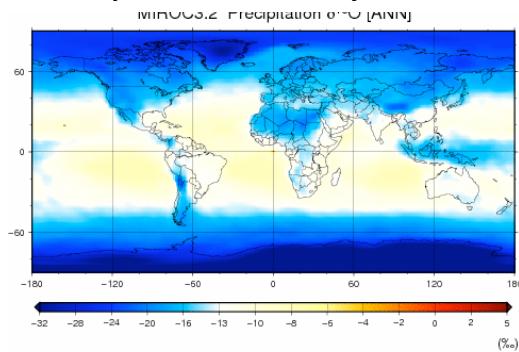
Surface Isotopic distribution



Isotopic field of precipitation



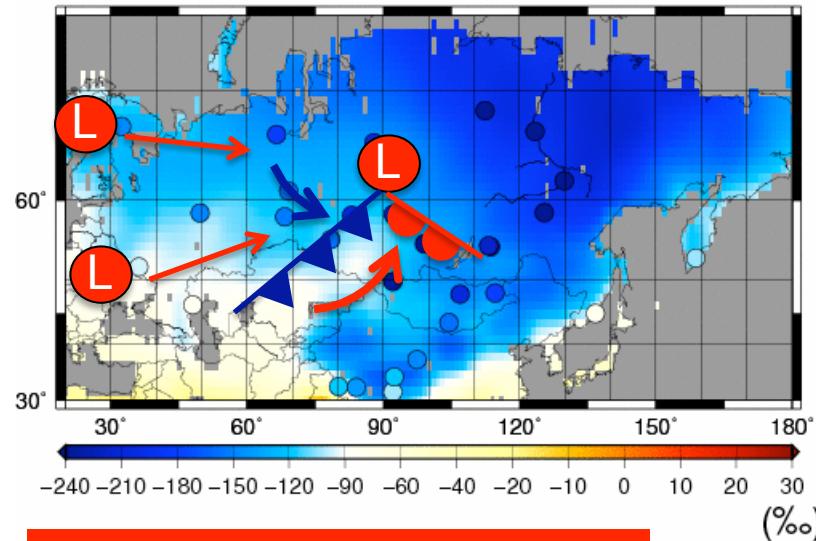
Isotopic variability due to rain-out



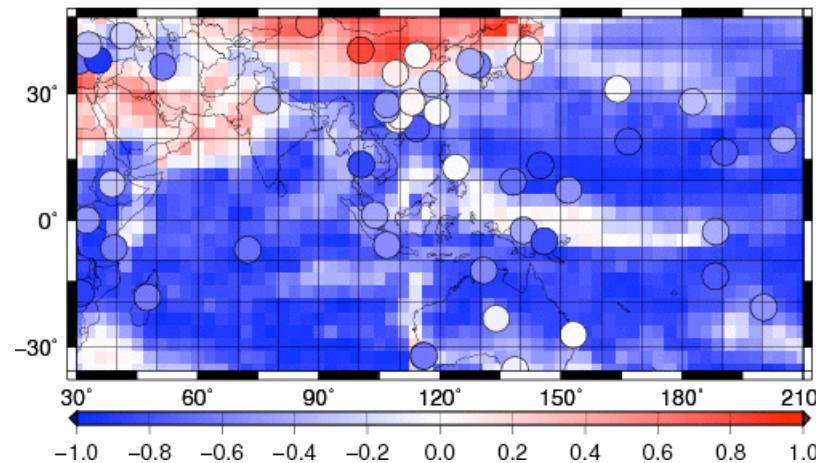
$$\delta(\text{prec}) = \delta(\text{source}) + \varepsilon \times \text{ROI} \quad (\varepsilon: \text{fractionation factor})$$

Key factor for isotopic variation

Middle & High latitude region



Tropics & subtropics region



Correlation between $\delta^{18}\text{O}$ and rainfall

Isotopic variability of precipitation is related with moisture transportation pattern associated with low pressure system.



Moisture source contribution

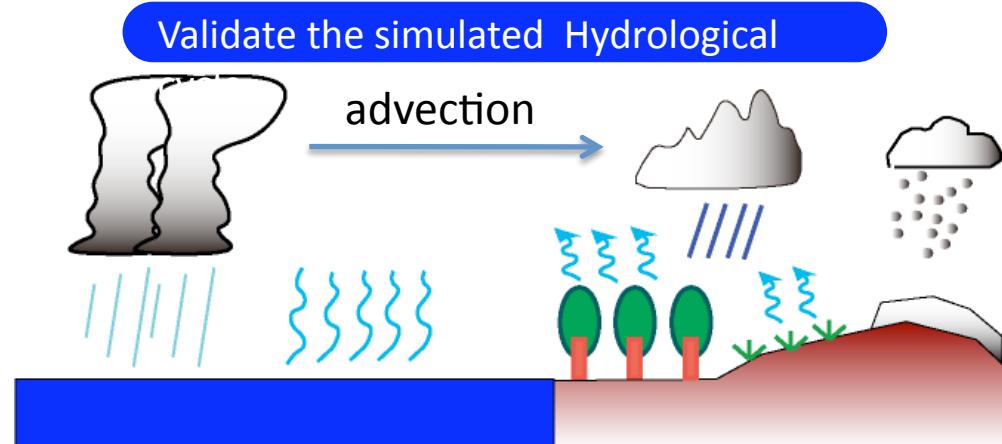
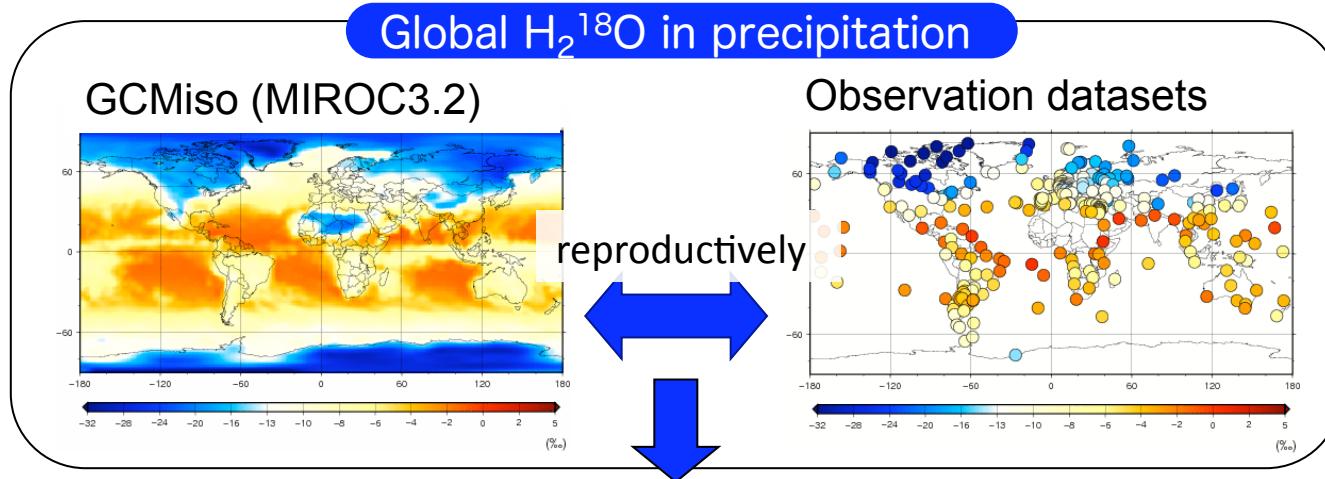
Isotopic variability of precipitation has negative correlation with rainfall amount.



Regional rain-out contribution

Diagnosis of the simulation

Isotopic variability in the atmospheric water is closely related with integrated history during transportation from source region. Thus, good reproducitivty of water isotope field results in the model can successfully simulate the hydrological cycle in the atmosphere.



A photograph of a polar bear standing on a large, white, textured iceberg. The bear is facing towards the left of the frame. The background consists of dark, choppy water under a clear sky.

Thank you for your attention!

September 2008@ Arctic Ocean